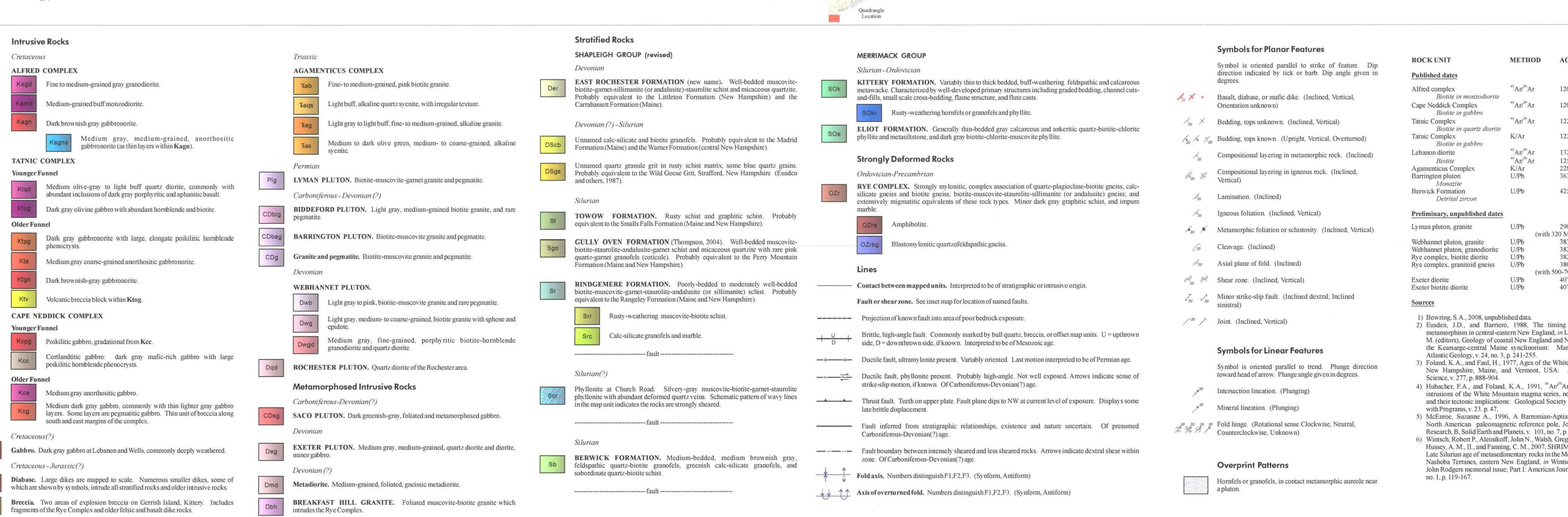
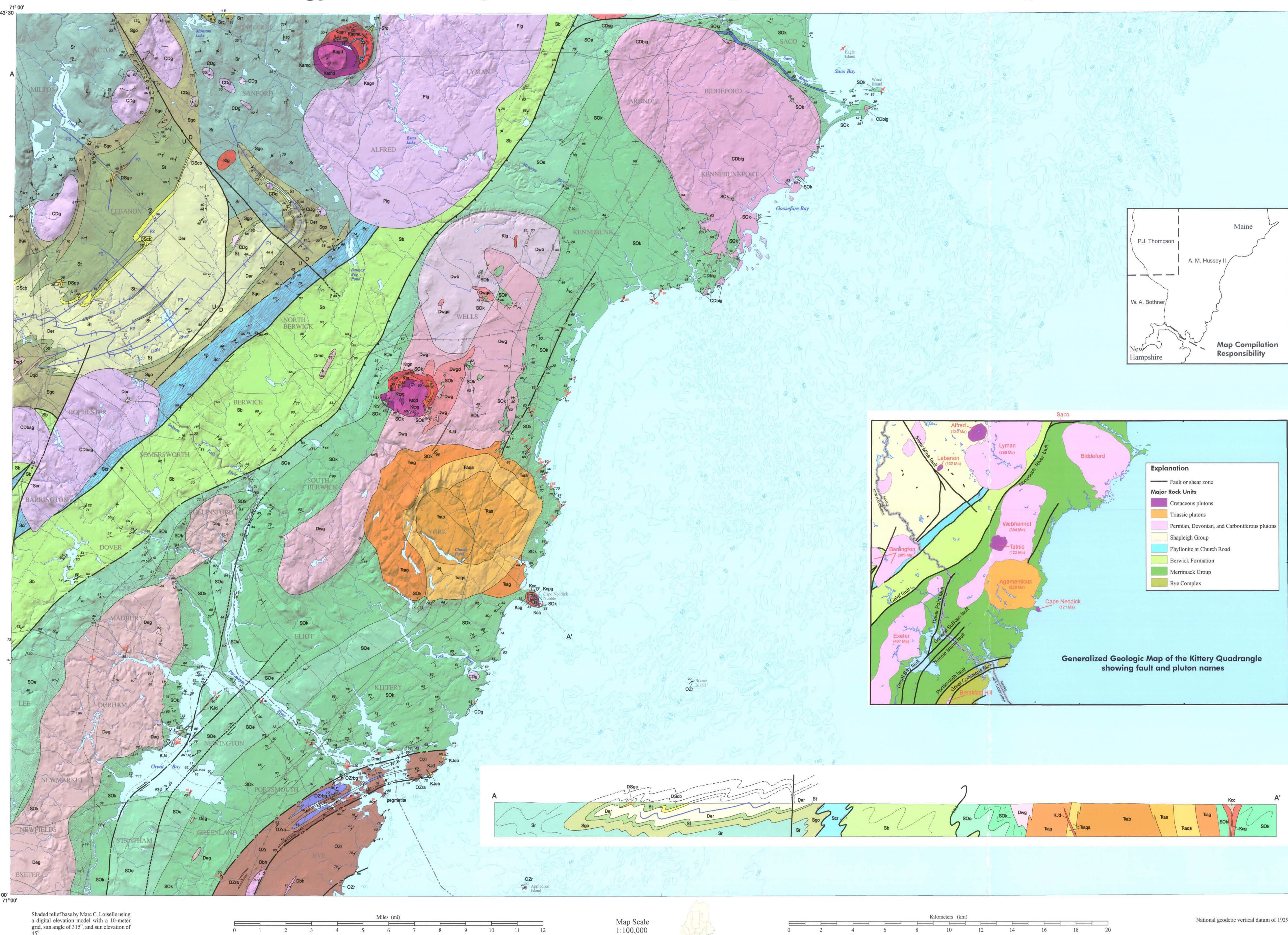


Bedrock Geology of the Kittery 1:100,000 Quadrangle, Maine and New Hampshire



Kittery Quadrangle

Maine-New Hampshire

Bedrock geologic mapping by
Arthur M. Hussey II, Wallace A. Bothner, and Peter J. Thompson

Digital cartography by:
Susan S. Tolman

Robert G. Marvinney
State Geologist

Cartographic design and editing by:
Robert D. Tucker

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Maine Geological Survey
Address: 22 State House Station, Augusta, Maine 04333
Telephone: 207-287-2801 E-mail: rgm@state.mn.us
Home page: <http://www.state.mn.us/doc/crimcrim.htm>

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STRATIFIED ROCKS

Stratified rocks in the Kittery quadrangle are metamorphosed deposits of shale, siltstone, and sandstone that accumulated in two ancient ocean basins, the Central Maine basin on the west and the Merrimack basin on the east. Deposition began during latest Ordovician through Middle Silurian time in the Merrimack basin, and from latest Ordovician through Early Devonian time in the Central Maine basin. These rocks were derived from weathering and erosion from exposed landmasses surrounding these basins. Sedimentary rocks are separated into formations of closely similar rock types that can be mapped over a broad area. Within the Kittery quadrangle these include the Eliot and Kittery Formations deposited in the Merrimack basin, and the Berwick Formation, the source material for the phyllonite at Church Road, and formations of the Shapleigh Group (Rindgemere, Gully Owen, Towow, and East Rochester Formations) accumulated in the Central Maine basin.

The Eliot Formation consists of metamorphosed thin beds of alternating tan-weathering limy siltstone, and dark gray shale (Photos 1 and 2). The Kittery Formation consists of variably thin to very thick-bedded metamorphosed limy sandstone, siltstone, and shale that preserve a variety of sedimentary structures similar to those found today in deep-sea sediments that accumulate at the base of continental slopes. Graded bedding (Photo 3) and flute casts (Photo 4) are common and enable us to determine stratigraphic order and the direction from which the sediments were derived and transported. In this case, from a continental land mass lying to the east (in terms of present day orientation) of the Merrimack basin.

The Berwick Formation consists of metamorphosed calcareous sandstone, siltstone, and shale much like the Kittery Formation but of more medium- to massive-bedded character (Photo 5). The Rindgemere, Gully Owen, and East Rochester Formations consist of metamorphosed gray non-rusty-weathering shale, siltstone, and sandstone. Bedding of the Rindgemere Formation is weakly developed and variable, and primary sedimentary structures are rare. On the contrary, bedding in the Gully Owen (Photo 6) and East Rochester (Photo 7) Formations is characteristically well developed, with rhythmic alternations of shale and siltstone which are commonly graded (lighter colored siltstone grading up into darker colored shale followed by an abrupt break to the overlying siltstone, Photo 6). The Gully Owen Formation includes distinctive thin beds of reddish granite-quartz granofels (Photo 8). The Towow Formation is a distinctive alternation of siltstone, easily weathering poorly bedded to massive metamorphosed shale and siltstone. Between the Towow and East Rochester Formations lies an unnamed thin, possibly discontinuous, unit of metamorphosed calcareous sandstone, and a thin rusty-weathering metasilstone or metacalcium with granule-sized clasts up to 4 mm in size consisting of white and blue quartz, quartzite and rare metacalcic chips.

INTRUSIVE IGNEOUS ROCKS

The major intrusive bodies (called plutons) of the Kittery sheet include the following:

- 1) Biddeford pluton, a medium-grained evenly textured gray biotite granite (Photo 9).
- 2) Webhannet pluton, consisting of three principal phases. The youngest, forming the northern end of the pluton, is a pink to light gray medium-grained evenly textured biotite-muscovite granitic; the intermediate-aged phase is medium- to coarse-grained, evenly textured to slightly foliated, light to medium gray biotite granite with moderate amounts of plagioclase and epidote; the oldest phase is fine- to medium-grained medium gray biotite-hornblende granodiorite.
- 3) Lyman pluton, a fine- to medium-grained, irregularly textured biotite and biotite-muscovite granitic. Regular masses of granitic pegmatite are common in many parts of the pluton.
- 4) Exeter pluton, medium-grained variably textured quartz-biotite diorite with minor irregular masses of comagmatic gabbro to quartz monzonite.
- 5) Barrington pluton, medium- to coarse-grained biotite diorite.
- 6) Minor unnamed plutons mostly of biotite-muscovite granite with feldspar phenocrysts up to 3 cm long, commonly aligned with the foliation, which forms a swirled pattern in many outcrops. Pegmatite is irregularly distributed, especially in the Rindgemere Formation, as isolated bodies or magmatic segregations.
- 7) Agamenticus Complex, consisting of four phases closely associated in space and time. The oldest is medium- to coarse-grained olive green syenite rocks similar to granite but with very little quartz. The next younger phase is buff quartz syenite with very irregular texture. The next phase is a light gray to slightly pinkish gray, fine- to medium-grained granitic. These three phases have a very unusual variety of dark minerals including riebeckite, arfvedsonite, enigmatite, actinolite, and ferrosillimanite which reflect the calcium-rich nature of the magma that gave rise to them. The youngest phase of the complex is fine- to medium-grained pinkish gray biotite granite.
- 8) Gabbro complexes. These are relatively small plutons (Cape Neddick, Tatic, Alfred, and Lebanon plutons) that consist of a variety of iron- and magnesium-rich igneous rocks such as gabbro, ceratophane, monodiorite, and granodiorite. These are the youngest crustal rocks in the Kittery quadrangle, having been emplaced about 120 million years ago.

In addition to these plutons, thin tabular igneous rock masses (dikes and sills) ranging from a few centimeters to 75 meters in thickness intrude all the metamorphosed stratified rocks and all the igneous plutons except the gabbro complexes. They are especially common in the exposures of the Kittery Formation in the coastal area from Kittery to Biddeford Pool (Photo 16). They are predominantly basaltic diorite (a coarse grained phase of basalt), approximately the same composition as gabbro.

DEFORMATION AND METAMORPHISM

The stratified rocks have been extensively folded and faulted. Two stages of folding are particularly well displayed in the Kittery Formation along the coast between Ogunquit and Kittery. Earlier folds are recumbent, that is, they are lying over on their side (Photo 10), whereas the later folds are upright, both sides of a fold dip steeply as illustrated in Photo 11. Major faults in the map sheet include the Nonesuch River fault separating the Berwick and Eliot Formations, the Portsmouth fault separating the Kittery Formation and Rye Complex, the Great Common fault cutting the Rye Complex, and a shear zone embayed in the phyllonite at Church Road (Photo 12) defining a boundary between the Berwick Formation and other members of the Central Maine basin. Old silver mines and prospect pits of the Acton Silver District are located along a zone of white bull quartz, apparently marking the trace of a Mesozoic fault (Silver Mine fault) in the Acton-Lebanon area.

Most of the large scale folding is probably the result of the Acadian orogeny, a major period of middle Paleozoic crustal compression driven by plate tectonic movements. An additional effect of this event was the recrystallization of the stratified rocks to form schists out of shales and siltstones, and granofels and gneisses from sandstones. The degree of heating during this compression is expressed in the variety of minerals that formed, particularly in the schists. Lower temperature, for example, is indicated by the presence of chlorite, medium temperature by garnet, and higher temperature by sillimanite, and andalusite.

GEOLOGICAL HISTORY

The oldest stratified rocks are those of the Rye Complex. They consist of layered metasedimentary rocks (Photo 13) that were once important sandstones, shales (some carbonaceous), impure limestones, and perhaps a few volcanic rocks (Photo 14). Affected by at least two episodes of Paleozoic deformation, metamorphism (including partial melting), and intrusion, followed by suchthrusting but variable, shear (Photo 15), they no longer preserve the primary structures that hint at source area. The major basin in which these rocks accumulated is unclear since their separation from rocks of the Kittery and Eliot Formations by the Portsmouth fault.

The Eliot and Kittery Formations were deposited during Late Ordovician to Early Silurian time along the eastern edge of a deep seaway, the Merrimack basin, bordering an eastern landmass known as Avalonia. Primary sedimentary structures such as those shown in Photos 2, 3, and 4, indicate that sediments of the Kittery Formation came from an easterly source. Deposition in the Merrimack basin ceased by Middle Silurian time, followed thereafter by deformation and metamorphism related to an early phase of the Acadian orogeny. The Berwick, Rindgemere, Gully Owen, and Towow Formations accumulated in a western part of the seaway (the Central Maine basin), with sediments derived from erosion of the Bronson Hill terrane, an island arc complex of Ordovician age. However, by the time of deposition of the East Rochester Formation the source area shifted to the east, probably the uplifted and deformed rocks of the Merrimack basin. Deformation of the rocks of the Central Maine basin took place in early Devonian time as a result of a later phase of the Acadian orogeny.

Magmatic activity spans the middle Paleozoic to middle Mesozoic. The older intrusive rocks (e.g., 407 Ma Exeter pluton, and undated Rochester diorite) cut early Acadian folds in the southeast area of New Hampshire and Maine. Granitic and granodioritic rocks of the Webhannet, Biddeford, Barrington and North Lebanon and related small plutons were intruded at about 380 Ma, close to the end of the Acadian deformation. Granitic intrusion continued into the late Paleozoic (~260 Ma) and is represented here by the Lyman pluton. It is likely related to the Sebago batholith, just north of this map area, and one of the largest batholiths in the New England area. Some investigators relate the emplacement of these youngest Paleozoic granites to the Appalachian orogeny, the terminal compressional event of the Paleozoic Era in the Appalachian Mountains.

The last major event in the geologic history represented on the Kittery sheet is the intrusion of rocks of the White Mountain Plutonic-Volcanic Series of Mesozoic age. The Agamenticus Complex is a ring complex of Triassic age (235 Ma), one of the oldest intrusive rock bodies of the series related to the opening of the Atlantic Ocean. It is part of a north-south-trending chain of related volcanic centers from the New England area, extending across central New England and southern Quebec. This rift and rift-related magmatism spanned over 100 Ma. The last vestiges of this age are reflected in the ~120 Ma age gabbro complexes exposed in Tatic Hill, Alfred, Cape Neddick, and Lebanon. In addition to these complexities, the area is crisscrossed by abundant basalt and diabase dikes (Photo 16), some of them tens of meters thick and traceable for tens of kilometers across the map area, and hundreds of kilometers to the northeast and southwest. They maintain an orientation roughly parallel to edge of the North American plate where it splits and continues to separate from the Eurasian plate. A majority of these dikes are of Jurassic age (greater than 145 Ma and less than 205 Ma), emplaced after the Agamenticus Complex but before the gabbroic complexes.

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Geologic Time Scale

Geologic Age	Geologic Time Scale	Absolute Age*
Cenozoic Era	Cenozoic Period	65-145
	Quaternary Period	125,000 to 1,500,000 (0.02 to 1.5 Ma)
Mesozoic Era	Cretaceous Period	65-145
	Deonian Period	145-200
	Triassic Period	200-252
	Permian Period	252-300
	Carboniferous Period	300-360
Paleozoic Era	Devonian Period	360-418
	Silurian Period	418-444
	Ordovician Period	444-489
Precambrian	Cambrian Period	489-541
	Other than 544	541-571

* In millions of years before present. (Oriskany, A. V., 2002. Eché des temps géologiques, 2002. Université de la Nouvelle-France, Québec. <http://www.univ-nouvelle-france.ca/ressources/la-terre-ancienne/geologie/REVISION3081>.)

Photo 1. Thin, folded beds of the Eliot Formation at the type locality at Lost Fox Corner, New Hampshire.

Photo 2. Glacially polished Eliot Formation, gravel pit in Saco, Maine.

Photo 3. Graded beds in the Kittery Formation along the shoreline in York, Maine. A graded bed (two of which are well shown) has a light gray metacalcium lower part that grades upward into a dark metacalcium, the top of which is an abrupt break to the overlying metacalcium. These beds, top to the right of the photo and are inclined because of folding that has affected these rocks. The conspicuous graded bed in the center of the photo is approximately eight centimeters thick.

Photo 4. Flute casts in Kittery Formation at Moody Point, Wells, Maine.

Photo 5. Berwick Formation at type location, South Berwick, Maine, Salmon Falls River. Note typical thin to thick bedded character of the formation.

Photo 6. Bedding typical of the Gully Owen Formation at an exposure along Goding Road, Sanford, Maine. Note the conspicuous graded bed under the notebook (light colored metacalcium grading up into dark colored metacalcium). These beds are steeply inclined because of folding during the Acadian orogeny.

Photo 7. Bedding typical of the East Rochester Formation exposed in Little River, Lebanon, Maine. Arrow points to a graded bed, indicating that these beds are overturned.

Photo 8. Thin, graded quartz granofels (conglomerate) bed in the Gully Owen Formation, Route 160 cut, Rochester, New Hampshire.

Photo 9. Strangers of the Biddeford granite (light colored and massive) cutting the Kittery Formation (darker), Biddeford Pool, Maine.

Photo 10. Earlier recumbent folds of the Kittery Formation, Marginal Way, Ogunquit, Maine.

Photo 11. Upright folds of the Kittery Formation along York, Maine, shoreline.

Photo 12. Phyllonite at Church Road along Turkey Street, North Berwick, Maine.

Photo 13. Sheared beds of metacalcium (medium gray) and metacalcic phyllonite of typical of the Rye Complex at Fort Foster, Kittery, Maine.

Photo 14. Thinly laminated amphibolite, possibly representing mafic volcanic rocks within the Rye Complex at Seward's Point, Kittery, Maine.

Photo 16. A 2-meter-thick diabase dike cutting the Kittery Formation, in turn cut by a thinner basaltic dike. Shoreline north of Bald Head Cliff, York, Maine.