

## **Companion Report to accompany**

### **Bedrock Geologic Map of the Lake Francis 7.5' Quadrangle, New Hampshire**

**David R. Converse, Wallace A. Bothner, Christian E. Jahrling, II, and  
Philip S. Koch<sup>1</sup>**

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**Shane Csiki, New Hampshire State Geologist  
New Hampshire Geological Survey, 29 Hazen Drive, P.O. Box 95, Concord, NH  
03302**

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<sup>1</sup> David R. Converse ([drconverse7@gmail.com](mailto:drconverse7@gmail.com)), 2119 McGraw Ranch Road, Estes Park, CO 80517; Wallace A. Bothner ([Wally.Bothner@unh.edu](mailto:Wally.Bothner@unh.edu)), Professor Emeritus, Earth Sciences, University of New Hampshire, James Hall Rm 334, Durham, NH 03824; Christian E. Jahrling, II ([cjahrling@outlook.com](mailto:cjahrling@outlook.com)), 4240 NW 107th Avenue, Apartment 4401, Doral, FL 33178, Philip S. Koch ([pskoch56@icloud.com](mailto:pskoch56@icloud.com)) 14406 Wamblee Trail, Conifer, CO 80433.

## **Introduction**

This area is sparsely populated by year-round residents, but with many part-time residents and visitors. It is a favorite location for recreation – both winter and summer/fall activities. Forestry has been a major industry in the region for many years.

The first mention of its geology is found in Charles Hitchcock's 1877 Geology of New Hampshire and his map folios. Subsequent mapping was performed nearby by Billings (1956), his students (Hatch, 1963; Green, 1964, 1968; Myers, 1964), and later in the quadrangle by Converse (1977), Jahrling (1983) and Bothner, Jahrling and Moench (unpublished mapping 1980 – 1990). The USGS published a regional map (2° sheet) in 1995 (Moench et al., 1995 and references therein) and a modified version of that regional map was included in the Geologic Map of New Hampshire (Lyons et al., 1997). The bedrock geology of northeastern Vermont was compiled by Ratcliffe et al. (2011). Geophysical maps are available in Bothner et al. (1997).

The bedrock geology of southern Quebec that is adjacent to the New Hampshire border was mapped in the last decade by Tremblay et al. (2015) and Perrot (2019).

The aim of this project is to provide modern bedrock geological maps of northernmost New Hampshire which can be used for both practical and scientific applications. The Lake Francis Geologic Map is the second of 5 geologic maps that will cover the 1926 Indian Stream 15' quadrangle. The Geologic Map of the Pittsburg 7.5' Quadrangle (Bothner et al., 2020) was completed in September, 2020. Mapping in the Cowen Hill Quadrangle is already underway. Further refinement of this map is anticipated as the remaining maps in Northern New Hampshire are completed.

### ***Regional Bedrock Geologic Setting***

Northernmost New Hampshire is underlain largely by low grade metamorphic Silurian and Devonian rocks of the Connecticut Valley - Gaspé Trough (CVGT) as shown in Figure 1. The CVGT is fault-bounded on the west by Ordovician and older rocks of Taconic and Grenville orogenic belts and on the east by the Bronson Hill – Boundary Mountain belt (BHA). The CVGT recorded initial extensional tectonics as well as compressional tectonics as island arcs and microcontinents collided with

North America during the Siluro-Devonian Salinic Orogeny and the Late Devonian Acadian Orogeny.

The late Paleozoic Alleghanian orogenic and Mesozoic Atlantic rifting events are poorly represented at this latitude with the exception of plutons and dikes of White Mountain affinity.

### ***Geology of the Lake Francis Quadrangle***

#### **Surficial Geology**

This mapping effort did not focus on the surficial geology of the Lake Francis Quadrangle. The surficial geology consists of Quaternary sediments of fluvial and glacial origin. Glacio-fluvial gravel accumulations are found along streams in the area as well as on elevated terraces above current stream or river beds and are frequently mined for road materials. Glacial striations throughout the quadrangle record a consistent transport direction of  $\sim 140^\circ$  (S40°E). Large erratics are common particularly along Carr Ridge, which appears to be a lateral moraine.

#### **Bedrock Geologic Summary**

The bedrock geology consists of a few unmetamorphosed Mesozoic felsic and mafic dikes and of older Silurian and Devonian bedrock. The Silurian and Devonian bedrock consists of low-grade metamorphic slates, siltstones, sandstones and volcanoclastic sediments as well as metamorphosed intrusive and extrusive rocks. The mafic extrusive rocks are comprised of metabasalts and metabasaltic andesites, which were extruded as lava flows, pillow lavas, hyaloclastites, and lapilli tuffs. The felsic extrusive rocks are comprised of metadacites and metarhyolites, which consist of flows and tephra deposits (lapilli tuffs and ignimbrites). The intrusive rocks consist of diabase dikes and sills, and granitic sills and a small plutonic body.

Structurally, at least two major deformational episodes are recognized: 1) the upper Silurian-Lower-Devonian Salinic Orogeny (with both extensional and compressional events) and 2) the Late Devonian Acadian Orogeny compressional event. These deformations folded the older bedrock first in isoclinal folds and then refolded the bedrock into more open folds (e.g., Perrot, 2019, Perrot et al., 2018).

This map includes field data collected from 1976 through 2021. The field data includes structural, mineralogical and lithological data, new geochemical analyses, and five new age determinations. Data from surface exposures constrain the cross-section that shows the interpreted distribution of rock units in the subsurface. Key photos are also included to provide examples of different rock types, geometries and textures.

### **Geochronology**

No fossils have been discovered in the Lake Francis Quadrangle. An Emsian? fossil was discovered to the west in the Pittsburg Quadrangle (Hueber et al., 1990).

Crystallization and detrital zircon U-Pb age assignments were determined for meta-igneous and meta-sedimentary rocks in the area using data from GeoSep Laboratories, the IsoPlotR program and analytical methods described by Vermeesch (2021a, 2021b). Five new determinations were made in the Lake Francis Quadrangle as part of this work. Additional ages were available from the Pittsburg Quadrangle (Bothner et al., 2020), the southern Quebec (Perrot, 2019, Perrot et al., 2017, 2018, 2020) and the Second Lake Quadrangle (Dorais et al., 2017). Zircon age determinations are summarized in Table 1.

U-Pb age determinations on whole rock samples were made in the Second Connecticut Lake on a sample from the East Inlet Pluton by Lyons et al. (1986), and from a sample on Round Top Mt in the Lake Francis Quadrangle by Moench et al. (1995).

### ***Crystallization ages: Table 1***

Two samples: IS2020-151RV (biotite metagranite sill) and IS2020-94 (meta-ignimbrite) were analyzed and both yielded approximately the same crystallization ages:  $411 \pm 7$  ( $2\sigma$ ) Ma and  $411 \pm 3$  ( $2\sigma$ ) Ma respectively.

These ages are significantly younger than whole rock age dates from the East Inlet Pluton ( $430 \pm 4$  Ma – Lyons et al., 1986) in the Second Lake Quadrangle and from Round Mt in the Lake Francis Quadrangle (Moench et al., 1995). These ages are also younger than the zircon crystal age dates from felsic metavolcanics on the southern shore of Second Lake ( $432 \pm 8$  Ma – Dorais et al., 2017).

The older ages from the whole rock age dates are not surprising as metavolcanic samples from the Lake Francis Quadrangle contain significant zircon populations with older inherited ages, which likely would skew a whole rock age determination.

### ***Detrital Zircon ages: Table 1, Figure 2A & B***

Age date determinations from detrital zircons are typically used for two purposes:

- Establishing the maximum depositional age of the sample. The actual age of the sample can be younger, if the younger sediment did not contribute sufficient numbers or appropriate sizes of zircon crystals. For example, we noted that some of the pelitic metasediments in our area contained few and typically < 10 micron zircon crystals.
- Establishing sedimentary provenances with clusters of zircon ages indicating derivation from different sedimentary sources.

Three samples in the Lake Francis Quadrangle contained zircon crystals of multiple ages that were analyzed as “detrital” zircon samples.

- IS2020-112 was a metavolcaniclastic unit that contained lapilli that were typically < 3 cm. The maximum depositional age was determined to be  $408 \pm 6$  ( $2\sigma$ ) Ma.
- IS2020-207A – this sample had very small zircon crystals and very few young zircon crystals with high discordances yield an maximum depositional age of  $443 \pm 16$  ( $2\sigma$ ) Ma. Removing the high discordancy samples, yielded a maximum depositional age of  $603 \pm 17$  ( $2\sigma$ ) Ma
- IS2019-242 – only recovered 13 zircons from this sample, very few young zircons, maximum depositional age of  $459 \pm 16$  ( $2\sigma$ ) Ma.

### ***Stratigraphy***

We recognized three formations in the quadrangle –the Ironbound Mountain (tentatively), the Gile Mountain and the Frontenac Formation. The latter is separated into four members (3 meta-igneous units and 1 metasedimentary). A correlation chart showing the overall age relationships is shown in Figure 2A. Age dates from the Lake Francis Quadrangle indicate that the Frontenac Formation is approximately co-eval with the Gile Mountain Formation. However the sedimentary source terranes were different for the Frontenac and the Gile Mountain Formations (Figure 2B).

Many intrusive bodies (primarily metadiabase dikes or sills, and biotite granite) are mapped largely within the Frontenac and are described in detail in the Summary of Map Units on the map itself. The metamorphosed biotite granite may have been the source for the felsic metavolcanic rocks and ignimbrites as it has approximately same age and geochemistry.

Figures 3 and 4 show characteristic metasediments and metavolcaniclastics for the Gile Mountain and Frontenac Formations respectively. Figure 5 shows examples of pillowed metabasalts and a metadiabase intrusive. Figure 6 shows examples of metavolcaniclastic /meta-lapilli tuff unit from the north shore of Lake Francis, and of a metaignimbrite from Cedar Stream. Figure 7 shows an example of the metamorphosed biotite granite plutonic body on Tromley Hill with an inset of a thin-section with myrmekitic features from a similar outcrop on Round Top Mt, and an example of thin biotite granitic sills within a metabasalt on the south shore of First Connecticut Lake.

### ***Structure***

Broadly speaking, the phyllites and siltstones of Gile Mountain Formation exposed on the eastern flank of the shallow northeast plunging Beaver Brook anticline (the axis is in the Pittsburg Quadrangle) dominate the western portion of the quadrangle. This eastern limb is truncated by the west-directed Monroe Fault located east of the Lake Francis Dam.

Three faults are proposed in the map area – faults are very rarely exposed in outcrop in northern New Hampshire – the Lake Francis Quadrangle is no exception. Two of the faults, the Monroe and Deadwater Ridge Faults are shown on the Bedrock Geologic Map of New Hampshire (Lyons et al., 1997). The third tentatively proposed fault (here called the Cedar Stream Fault) is a normal fault that separates the Frontenac from a possible Ironbound Mountain equivalent in the southeastern corner of the quadrangle. All three faults were folded during the Acadian Orogeny.

The Monroe Fault separates the Gile Mountain Formation to the west from the Frontenac Formation to the east. Although ages of the two formations are quite similar (based on zircon age dating), the lithologies are distinct both in terms of volcaniclastic input, metaigneous units and the different sourcing based on zircon age distributions (Figure 2B). The second proposed fault is the Deadwater Ridge

Fault, another high-angle reverse fault separating the Frontenac metasediments to the west from the Frontenac metabasalts to the east. Both the Monroe Fault and the Deadwater Ridge Faults are consistent with geophysical modeling (gravity and magnetic) of the subsurface geometry (Jahrling, 1983). The proposed Cedar Stream fault is a normal fault (mapped as the Deer Pond Fault on the NH Geological Map (Lyons et al., 1997)) based on the inferred existence of the Ironbound Mountain Formation in the southeastern corner of the Lake Francis Quadrangle. The Ironbound Mountain Formation correlation is based on a lithologic comparison made over the years (e.g., Lyons et al., 1997), but is not supported by any age data. Our attempt in 2021 to date these rocks was unsuccessful due to the lack of zircon grains.

Abundant minor folds were identified that may be related to the formation of the Beaver Brook Anticline/Syncline pair in the Pittsburg Quadrangle but we also recognize the possibility of a more complex refolded nappe structure. These regional structures are present in both the Central Maine terrane and CVGT farther south in New England, often at higher metamorphic grade. We lack evidence of a regional inverted limb at this latitude.

Three deformation events, D1, D2, and D3 are recognized in the area. Rare isoclinal F1 folds, often with well-preserved graded beds (Pittsburg Quad, Cowen Hill), are refolded about tight, asymmetric generally westerly inclined, shallow northeasterly and southwesterly plunging mesoscopic F2 folds. Abundant F2 folds crop out along the northern shoreline of Lake Francis (Figure 8) at both large and small scales with low plunges (ca.  $10^\circ$ ). Late S3 cleavage in more pelitic layers sparingly represents D3.

### ***Frontenac and Gile Mountain Igneous Geochemistry***

In collaboration with Professor M.J. Dorais (BYU), we are investigating the whole rock geochemistry of the Frontenac metaigneous rocks to better understand the origin, tectonic setting and possibly develop a geochemically based igneous stratigraphy. Analytical results were received in early September. Figure 9 shows all data (separated into groups assigned by rock type and location) on two standard geochemical plots. Three preliminary results are: 1) metabasalts near the Monroe Fault appear chemically distinct from other metabasalts, 2) the metaignimbrite in Cedar Stream is nearly identical in composition with one of the metamorphosed

biotite granite samples, and 3) the meta-lapilli tuff sample from the north shore of Lake Francis has a rhyolite composition.

### ***Economic Geology***

As might be expected, in a submarine volcanic environment, there were a few signs of hydrothermal activity and associated mineralization. The most prominent example was a very small volcanogenic “massive” sulfide located on the southeastern shore of Lake Francis (Figure 10), with the word “massive” in quotes due to the very small size. The dominant sulfide is pyrite with some minor chalcopyrite as seen in a polished thin section from the sulfide-rich chlorite schist. The field and hand sample photos and an interpretive sketch map are shown in Figure 8. Due to the nearly vertical dips, the sketch map represents a cross-sectional view of the exhalative system.

### ***Acknowledgements***

We thank Josh Keeley of NHGS for his help with base map generation (Lidar and topographic), map layout and all things ARC-related and Rick Chorman, NH State Geologist and Shane Csiki (the new NH State Geologist), for continuing support of this STATEMAP project. We also benefited from a field visit with Bob Marvinney, State Geologist of Maine, and very helpful conversations with Morgann Perrot, Alain Tremblay as well as Lori Summa and David Awwiller. Paul O’Sullivan, GeoSep, Moscow, ID, analyzed zircons from metasedimentary and metaigneous rocks by LA-ICP-MS to determine U-Pb ages. Professor Michael Dorais (BYU) arranged for the geochemical analysis of the metaigneous rocks and provided helpful commentary. Chris and David Craigue kindly arranged lodging in Pittsburg, NH.

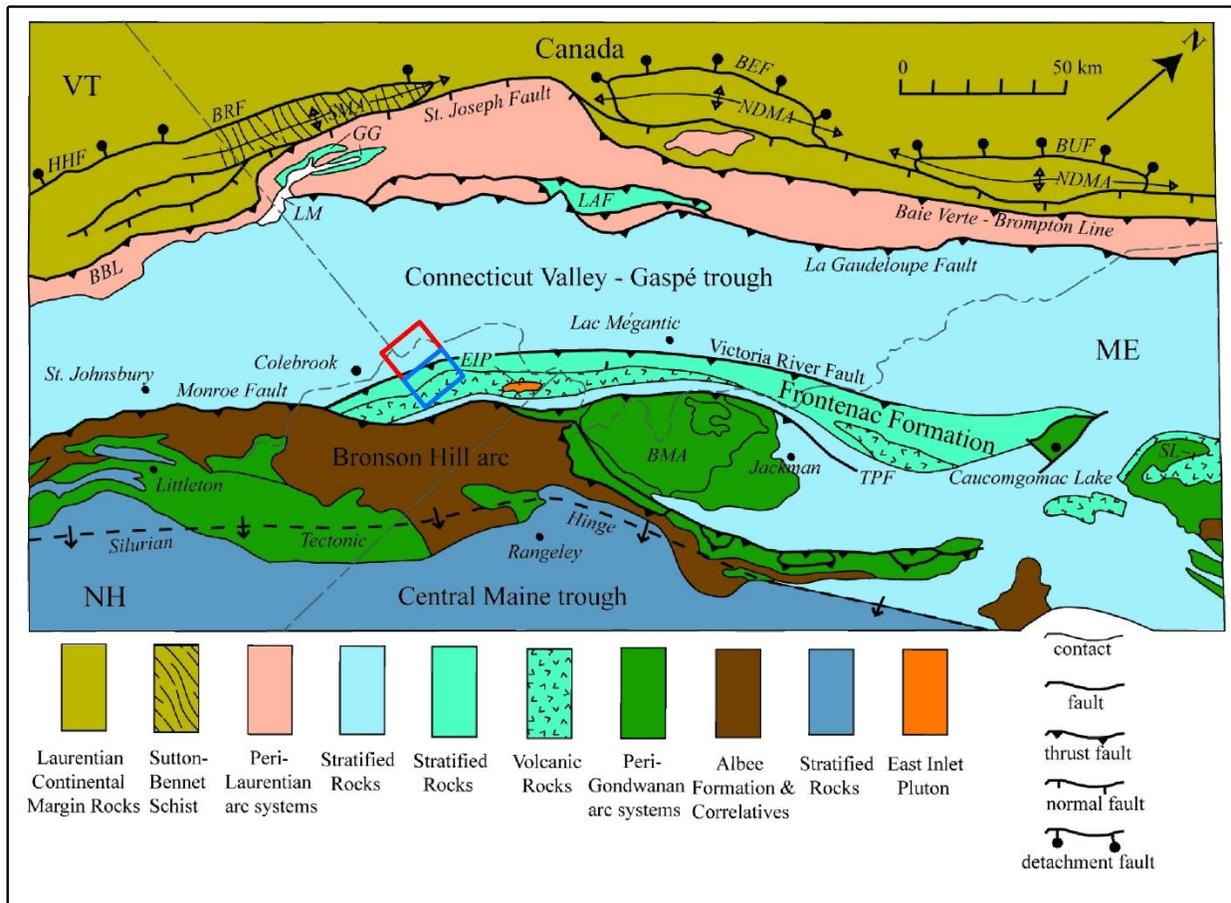
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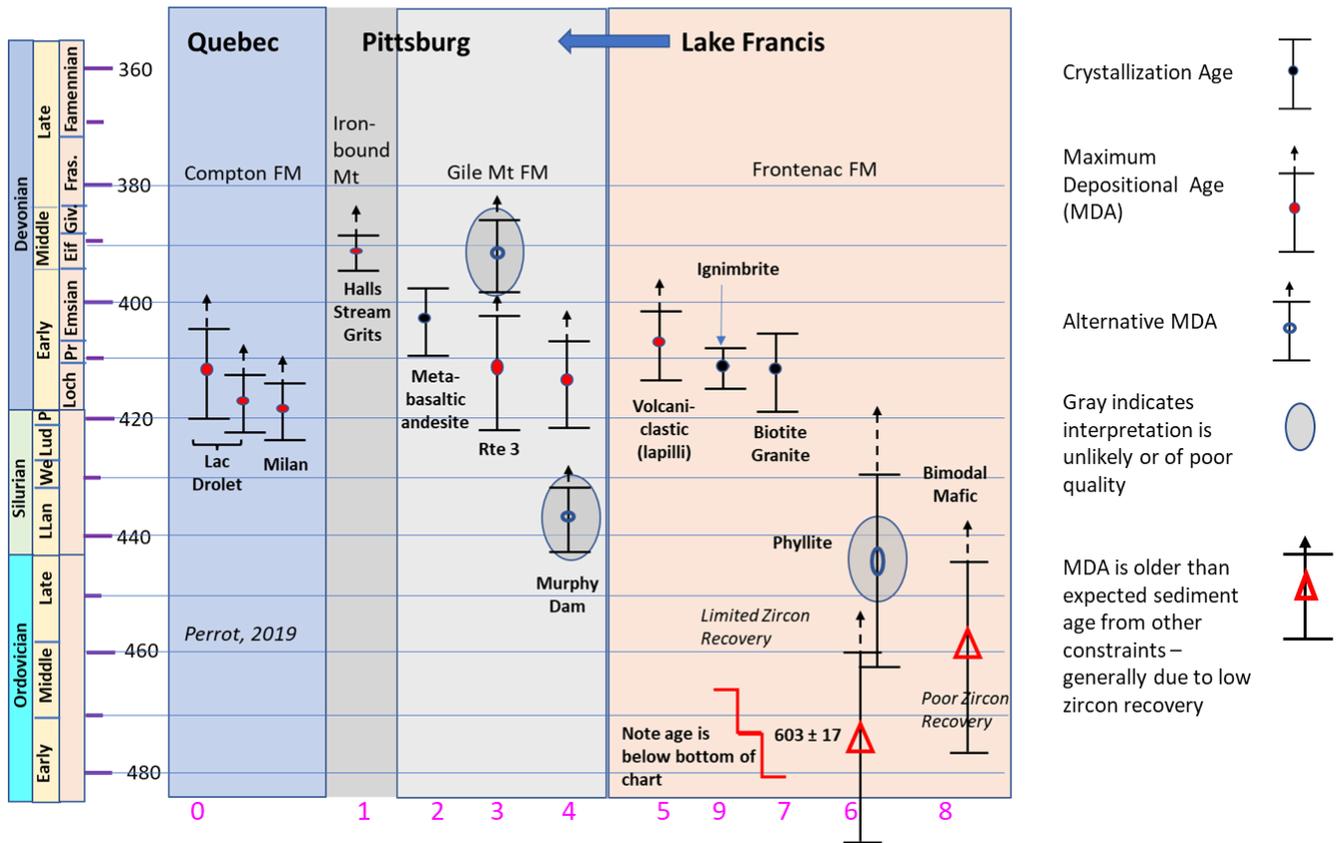
## Figures & Tables

**Figure 1 – Regional geology**

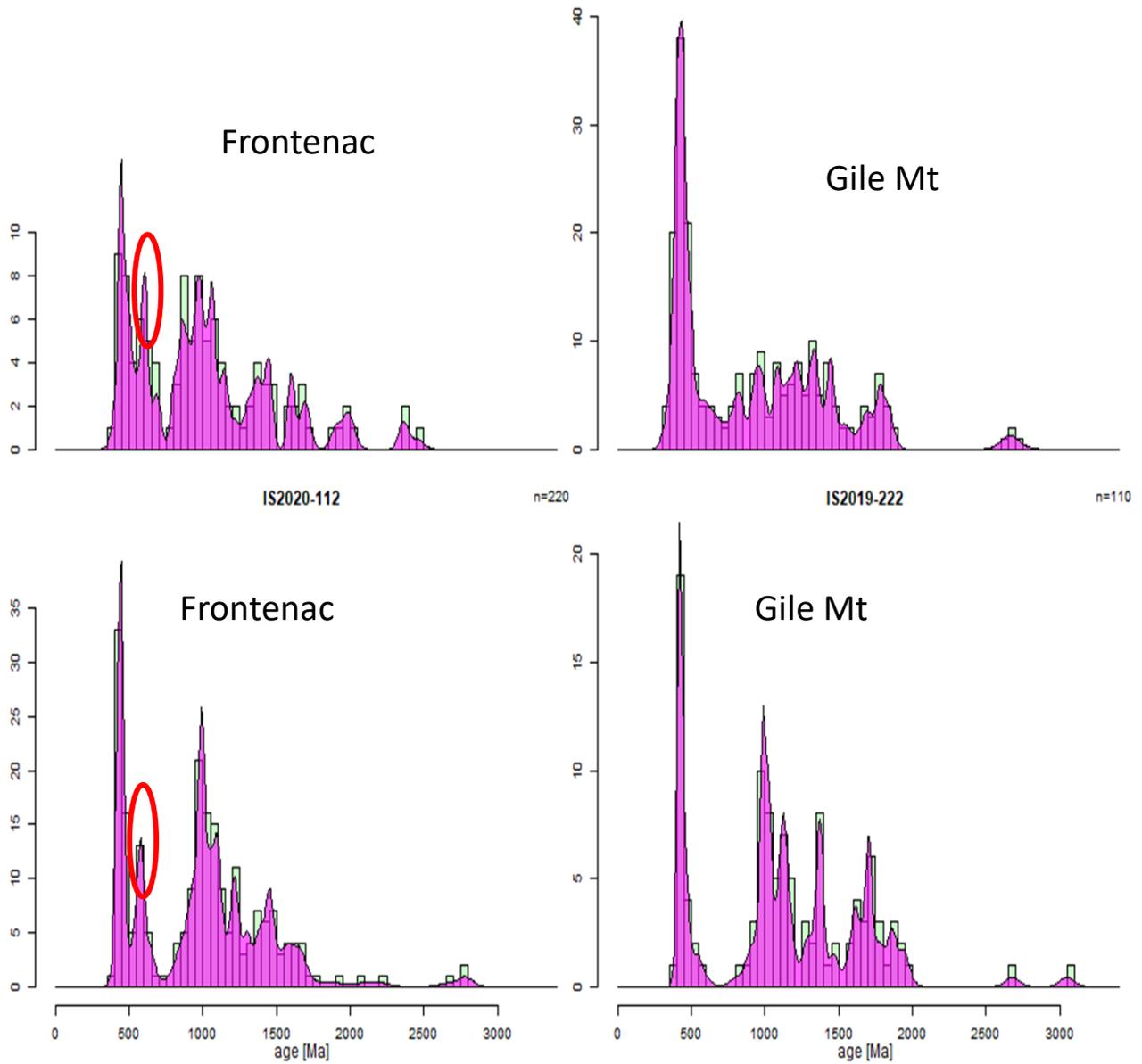


**Figure 1.** Simplified geologic map (modified after Dorais and others, 2017) showing the location of the Lake Francis Quadrangle (blue) and Pittsburg Quadrangle (red) at the southeastern border of the Connecticut Valley – Gaspé Trough. The Bronson Hill arch separates the CVGT from the Central Maine Trough. The Monroe fault is extended northerly to join the Victoria River (Belle) fault in Quebec and separates the Frontenac Formation from rocks of the CVGT in northern NH. EIP, East Inlet pluton; BMA, Boundary Mountains arch, BRF, Brome fault; BUF, Buckland fault; GG, Glenbrooke Group; HHF, Honey Hollow fault; LAF, Lac Aylmer Formation; LM, Lake Memphremagog; NDMA, Note Dame Mountains anticlinorium; SMA, Sutton Mountains anticlinorium; SL, Spider Lake; TPF, Thrasher Peaks fault

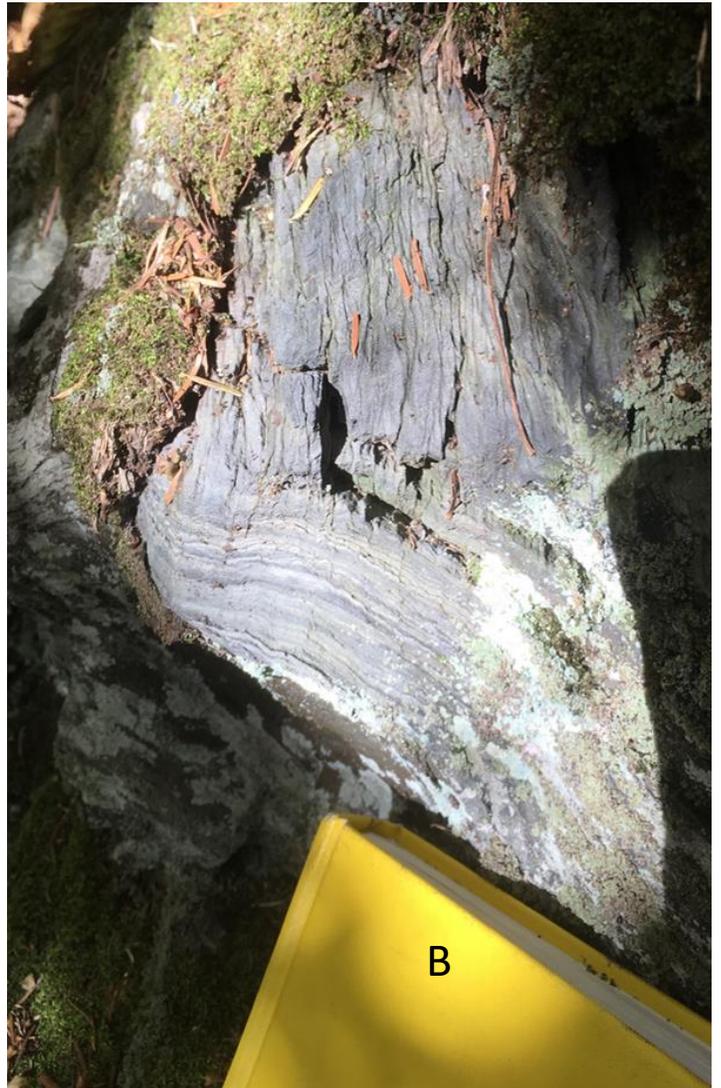
## Stratigraphic Age Constraints (Zircon U-Pb Age Dates)



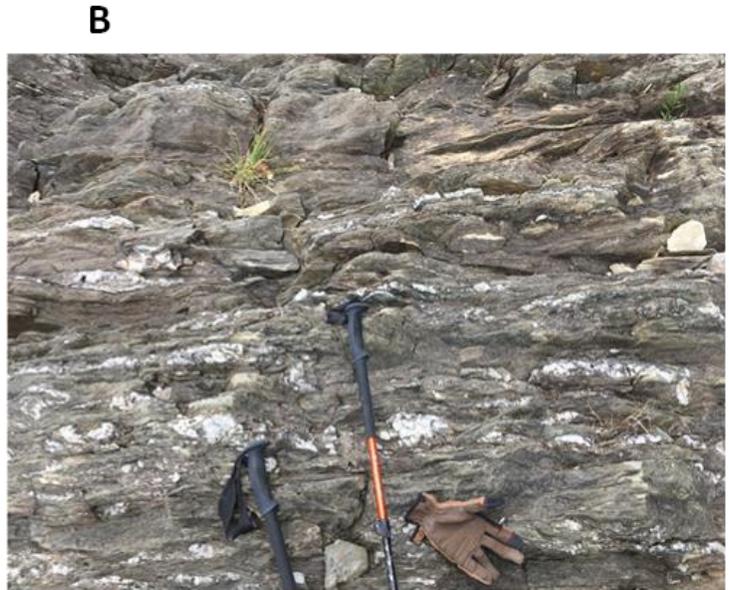
**Figure 2: A)** Zircon age determinations. Quebec dates are from M. Perrot (2018), the remainder are analyses of U-Pb from LA-ICPMS analysis conducted by GeoSep Laboratories. Two types of age determinations are shown – 1) crystallization ages and 2) maximum depositional ages (MDAs). It is important to understand that MDAs only provide an upper bound but not a lower bound for the depositional age – for two samples, the MDA is much older than the sediment age from other constraints, due to low zircon recovery). The age results are shown with 2σ uncertainty. In some cases, the data can support alternative interpretations for the MDAs – either due to limited number of zircons or issues with data quality. Note that a gray circle is added to indicate if an alternative interpretation is considered unlikely or of poor quality. Magenta numbers reference entries in Table 1.



**Figure 2: B)** Zircon age distribution in 2 Frontenac and 2 Gile Mt samples. The most obvious difference between the two formation is the lack of a ~ 582 Ma peak in the Gile Mt sandstones suggesting a different source provenance for the Frontenac Formation at that time.



**Figure 3:** A) Gile Mt metasilts and phyllites (IS2018-1) on southwestern shore of Lake Francis. Sedimentary features (scours) are complicated by the intense shearing nearby. B) Finely layered Gile Mountain gray phyllites and metasilts exhibiting partial turbidite sequences (IS2020-136).



**Figure 4:** **A)** Folded Frontenac volcanoclastic sediment on north shore of Lake Francis (IS2020-116), **B)** Frontenac Phyllite (IS2021-216A) along shore of 1<sup>st</sup> Connecticut Lake with abundant quartz veining. **C)** Thick-bedded Frontenac metasiltstone (Sherman Loop, IS2021-242) with bedding (S0) well displayed.

**Figure 5** : A) Frontenac Pillowed metabasalt (IS2020-215) crops out on the north shore of Lake Francis. Large pillow underneath the trekking poles clearly tops upwards (east), B) Frontenac Metadiabase (IS2020-213) photos showing metadiabase with interesting flow structures (in some cases pillows) and adjacent to the fine-grained flow structures is coarse-grained (> 3-5 mm) metadiabase indicating intrusion into a metabasalt pile.

**A**



**B**





**Figure 6: A)** Meta-Lapilli tuff /tuffaceous sediment on north shore of Lake Francis (IS2020-112, maximum depositional age  $408 \pm 6$  Ma) – geochemistry suggests a rhyolitic composition. **B)** Metaignimbrite outcrop in Cedar Stream (IS2020-94B). Geochemical composition is similar to metabiotite granite and zircon age is  $411 \pm 3$  Ma.



**A** Biotite granite cliff

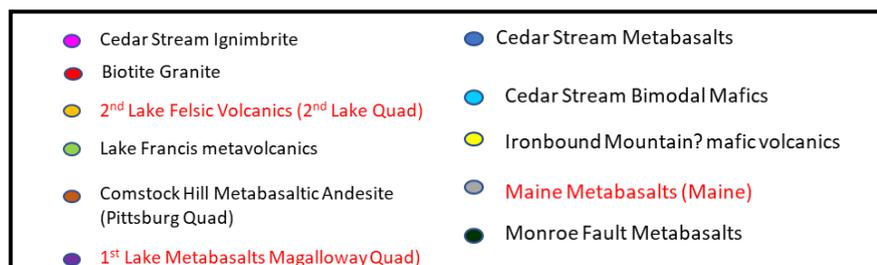
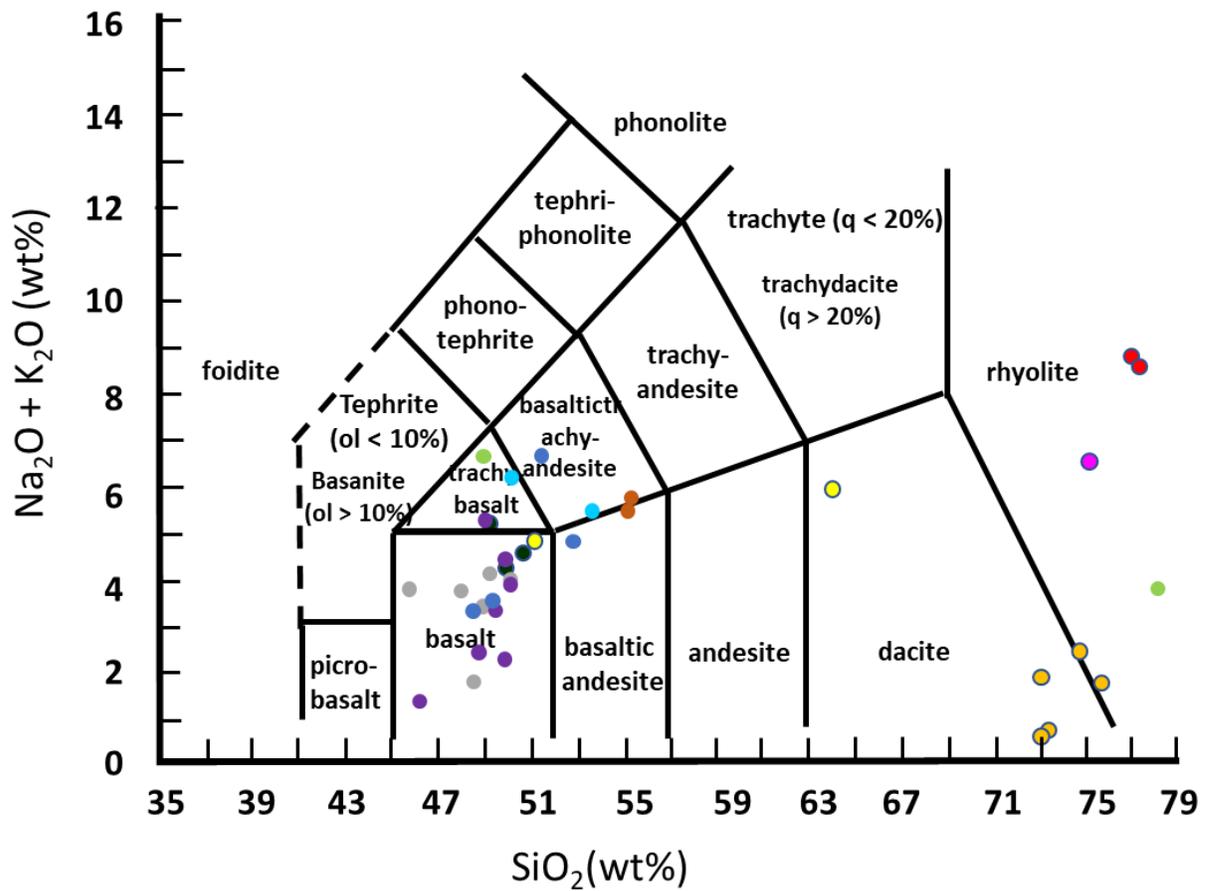


**B** Felsic sills in Frontenac metabasalt

**Figure 7:** **A)** Intrusive biotite granite forms large cliffs on Tromley Hill (IS2021-117). **B)** Biotite granite forms sills in the Frontenac metabasalt along the southern shore of the 1<sup>st</sup> Connecticut Lake (IS2021-17) and particularly in Cedar Stream / S Tromley Hill area (typically 1-3 m scale), where the contacts are knife sharp. Zircon crystallization age is  $411 \pm 7$  Ma. **C)** Thin section inset of “graphic” biotite granite with well-developed myrmekite from Round Top Mt (IS2020-144).

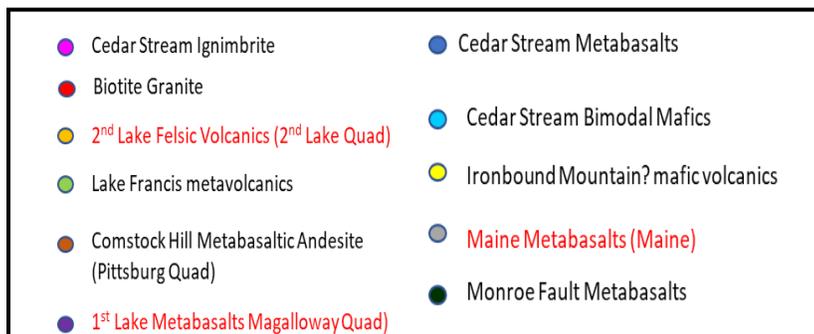
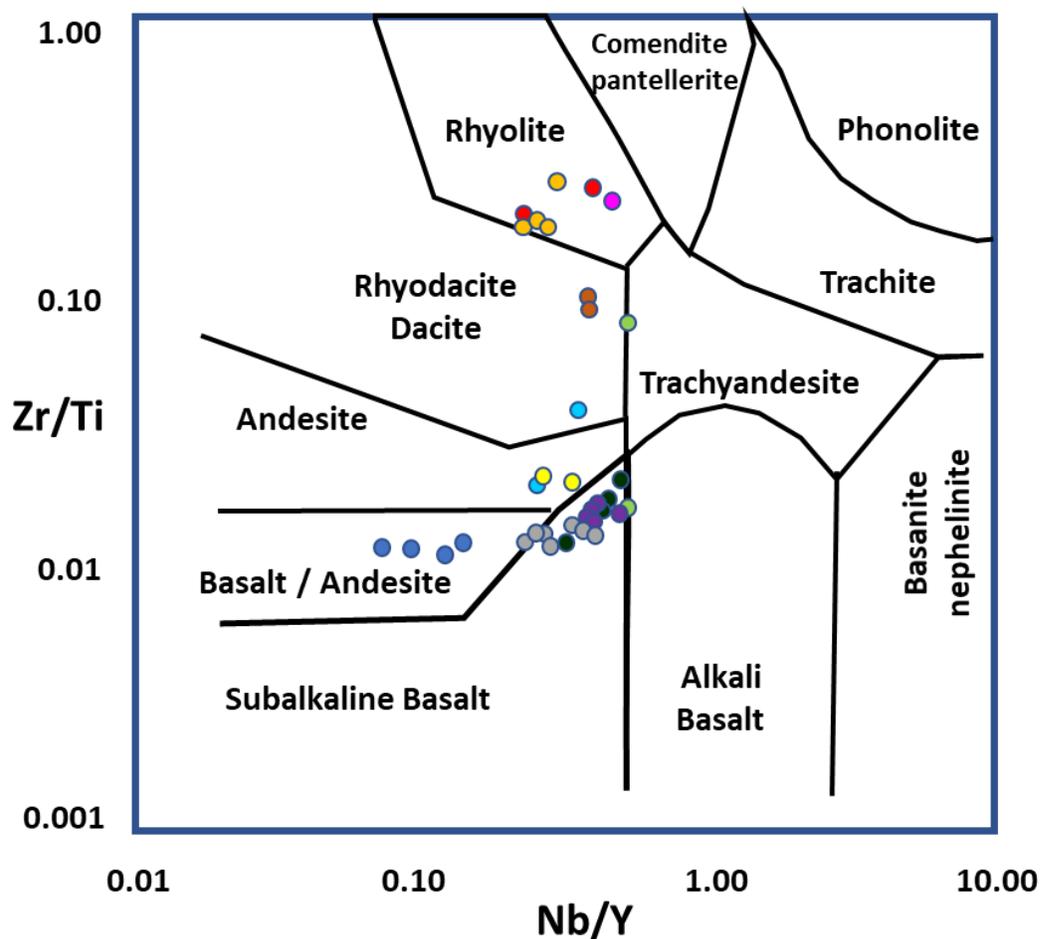


**Figure 8:** Large-scale open fold (F2) in Frontenac metavolcaniclastic sediment (IS2020-115). Note the quartz veining that outlines the fold. Folds of this size are rarely exposed in this area.



**Figure 9A: Frontenac Metalgneous Rocks and Gile Mt metabasaltic andesite**  
Major Element Chemistry – range of metabasalts, meta-basaltic andesites, dacites, rhyolites/granites.

*All red legend entries indicate data from Dorais et al., 2017; remaining data generated during collaboration with Professor M.J. Dorais (BYU)*



**Figure 9B: Frontenac Metagneous Rocks and Gile Mt metabasaltic andesite**

Trace element chemistry – see separation of metabasalts from Monroe Fault, similarity of biotite granite and metaignimbrite, unusual metabasaltic andesite from Gile Mt. (Pittsburg)

*All red legend entries indicate data from Dorais et al., 2017; remaining data generated during collaboration with Professor M.J. Dorais (BYU)*



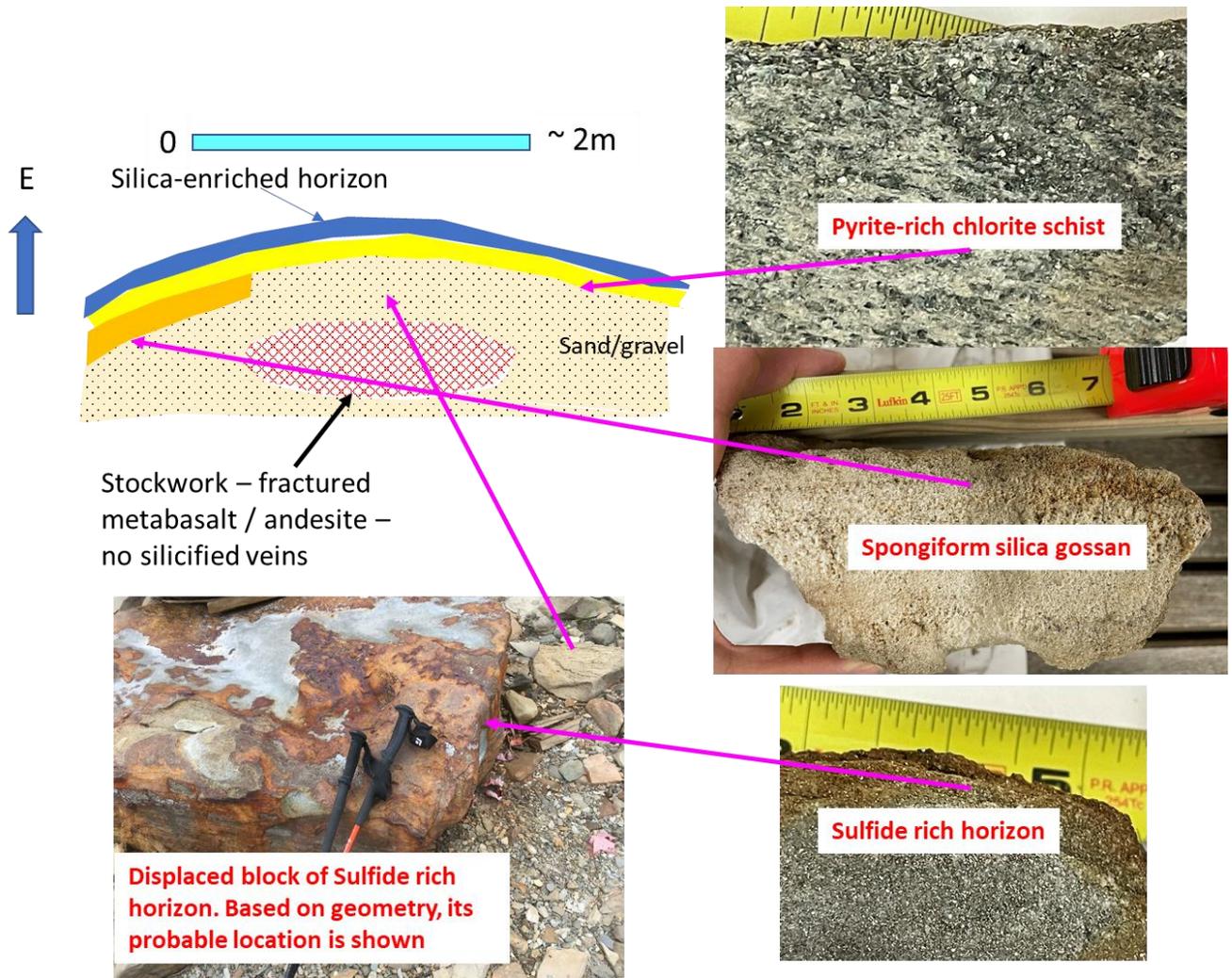
Sulfide-rich chlorite schist with abundant small pyrite cubes with minor chalcopyrite seen in polished thin-section



Large sulfide-rich block displaced ~ 5 meters away on the beach. Based on the geometry of the block, it may have come from the empty space as shown by magenta arrow.

Probable fractured stockwork zone. Did not see silica-rich veins filling the fractures.

**Figure 10A:** – Exhalative Deposit (IS2020-211): Small Volcanic “Massive” Sulfide deposit: exhalative sulfide deposit (partially dismembered).



**Figure 10B:** Small Volcanic “Massive” Sulfide deposit (IS2020-211): small exhalative sulfide deposit with hand samples shown from different zones. Sketch figure tries to capture the key elements of the exhalative system.

**Table 1: Summary of Zircon U-Pb Age Dates**

Sample	Location	Formation	Crystallization Age (Ma)	Maximum Depositional Age (Ma)	Comments
13MP26 0	Quebec near NH border	Lac Drolet		413± 7 (2σ)	From Perrot, 2019 Ph.D thesis
IS2019-172 1	Ridge W of Indian Stream (Pittsburg Quad)	Halls Stream Grits, Ironbound Mt		391 ± 2 (2σ) or 2 peak model 381-386 Ma and 397 -406 Ma	MWSD suggests two populations which may represent two events
IS2019-184 2	Indian Stream Comstock Hill (Pittsburg Quad)	Metabasaltic Andesite	403 ± 6 (2σ)		Good result
IS20-18 3	Tabor Notch Rd & US 3 (Pittsburg Quad)	Gile Mt		411 – 414 ± 8 - 10 (2σ) Alt: 391 ± 6 (2σ)	Loc. of Emsian fossils; Difficult younger ages
IS2019-222 4	Lake Francis Spillway (Pittsburg Quad)	Gile Mt		413 ± 8 (2σ) or 443 ± 5 (2σ)	Few young ages with discordance < 15%
IS2020-112 5	N Shore Lake Francis (Lake Francis Quad)	Front. meta-volcaniclastics		408 ± 6 (2σ)	Reasonable
IS2020-207A 6	NE Corner Lake Francis (Lake Francis Quad)	Front. meta-volcaniclastics		603 ± 17 (2σ) Ma Alt. 443 ± 16 (2σ) Ma	Very few “young” ages and high discordances.
IS2020-151RV 7	Cedar Stream (Lake Francis Quad)	Biotite meta-Granite	411 ± 7 (2σ)		Tight age distribution
IS2019-242 8	Cedar Stream (Lake Francis Quad)	Mafic metavolcanic		459± 16 (2σ)	Very limited younger age data.
IS2020-94 9	Cedar Stream (Lake Francis Quad)	Meta-Ignimbrite	411 ± 3 (2σ)		Tight age distribution
IS2020-102 10	SE corner Lake Francis (Lake Francis Quad)	Metabasalt		N/A	Did not recover any zircons
IS2020-56 11	Lang Hill, Lovering Mt. Quad	Calcium-rich Metadiorite		N/A	Only recovered 7 zircon grains